3D characteristics of snow bands and implications for surface snowfall in Northeast Winter Storms

Sandra Yuter & Laura Tomkins

Matthew Miller, Luke Allen, Kevin Burris, Declan Crowe, Jordan Fritz, Logan McLaurin, Toby Peele, the NASA IMPACTS Science Team, & Brian Colle (Stonybrook U.)



18th January 2024

NC STATE

UNIVERSITY

Terminology

- Use term "snow" for precipitation-sized ice particle that is large enough to fall in still air
 - Distinct from cloud-sized ice
 - "age of snow" time since snow particle first attained precipitation-size
- "microphysical pathway" sequence of d(mass)/dt changes a snow particle undergoes as a function of the sequence of environments it falls through along its path through the storm

Big picture takeaways from multi-year winter storm observations

Low correlation between enhanced Z in "snow bands" detected on regional scanning radar and hourly surface snow rates

- Usually observe coincident mixtures of snow particle shapes and degrees of riming
 - Mixtures yield varying distributions of shapes, sizes, and densities in the same volume which complicate interpretations and retrievals of snow rate
- In the 1-2 hours that it takes a precipitation-sized ice particle to fall from near cloud top to the surface, 3d ice streamers originating in generating cells are tilted and smeared
- Lack of vertical column continuity in local enhancements in radar Z



Hueholt et al. (2022)



Hueholt et al. (2022)



Ice growth forms

Ice growth form depends on air temperature **and** vapor content of air.

As particles fall and temperature and vapor content of environmental air changes, growth type can change.

Hueholt et al. (2022, BAMS)

Polycrystaline natural ice, Polycrystals ≠ Aggregates











PHIPS images courtesy of M. Schnaiter, KIT



*and local to ice edges by ventilation





Often get sequences of ice growth in different environments

Same color outlines two views of the same particle



Hueholt et al. (2022)

Sequential growth as particle falls to surface can yield complex 3D particles



NASA IMPACTS PHIPS images,

2 views of the same particle

initial columnar polycrystalline growth followed by tabular growth then branched growth



Lightly rimed



Graupel

Different degrees of riming

NASA IMPACTS field project PHIPS data (Schnaiter)

Graupel

Surface Observations from Multi-Angle Snowflake Camera (NCSU)

3 mm



Aggregationcollection of ice particles > 0.2mm diameter. yields jumbles of multiple individual ice particles called aggregates) Often includes ice particles with different shapes

2015 01 24 07 42 00, max diameter 17.8 mm, fall speed 1.33 m/s

Time-integrated state of individual snow particles --everything that happened to particle prior to observation

In this case,1) vapor deposition2) sublimation3) riming

500 µm ⊣



NASA IMPACTS field program PHIPS particle image (Schnaiter) and ER-2 Cloud Radar (McLinden, Li)

Time-integrated state of set of particles that end up in the same volume--Snow particles that initially formed separately often get mixed together

Surface Observations from Multi-Angle Snowflake Camera (NCSU)

500

Airborne PHIPS (Schnaiter)

μm

Particles with different shapes and degrees of riming often co-occur in the same volume



Airborne PHIPS (Schnaiter)

Surface

snowflake

camera (NCSU)

Mixtures of shapes and degree of riming based on airborne PHIPS ice particle images, minute by minute

P-3 Leg 2: 2/26/20 00:27 - 00:39



Near cloud-top flight leg sample for NASA P-3 closely coordinated with NASA ER-2

Processes that change reflectivity (Z)

In snow, changes in Z do not necessarily mean changes in mass per unit volume

Process	Change to IWC/LWC	Change to radar reflectivity
Riming	Increase	Increase
Vapor Deposition	Increase	Increase
Collision-Coalescence	Increase	Increase
Condensation	Increase	Increase
Aggregation	No change	Increase
Melting	No change	Increase
Evaporation	Decrease	Decrease
Sublimation	Decrease	Decrease
Freezing	No change	Decrease
Fragmentation	No change	Decrease
Raindrop Breakup	No change	Decrease



Mapping from volumetric water content to radar reflectivity has much larger uncertainty in snow compared to rain

	Rain	Snow
Density of individual particles	Constant	Varies with degree of riming and aggregation
Mass to shape relationship of individual particles	Well defined	<i>Varies</i> depending on microphysical pathway
Mass per unit volume to equivalent backscatter (shape, size, number)	Well defined if can assume exponential-like particle size distributions	<i>Varies</i> with different mixtures of snow with different shapes, riming, and aggregation properties

100 m scale aircraft observations from SNOWIE (Idaho) project



Figure adapted from Zaremba et al. (2023, JAMC)



Snow falls slowly ~1± 0.5 m/s, takes ~67 min to fall 4 km



Layers with snow have fair amount of time for integrated properties to accumulate

NASA ER-2 X-band radar (McLinden, Li) with VAD winds overlaid (Helms)

Commonly see wind speed and direction changes with height which can advect ice particles > 50 km in horizontal from origin generating cell



NASA ER-2 X-band radar (McLinden, Li) with VAD winds overlaid (Helms)

Illustration of vertical structures associated with horizontal Z features



NASA ER-2 radar (McLinden, Li)

Reflectivity [dBZ]



Winter storm

- mesoscale bands in radar reflectivity
 - Unlike warm season convective lines
- Some (esp. longer bands) are associated with strong frontogenesis
- Appear as transient features in a fluid rather than static entities moving with the mean flow
 - Better defined as local enhancements to dBZ field than as a fixed dBZ threshold

Ganetis et al (2018) analysis of data from 108 cool season storms (1996-2016). Strong frontogenesis increases the likelihood of a single band forming, but these long bands as well as shorter multibands also frequently occur in environments of weak frontogenesis to frontolysis



Distribution of 800-700 hPa Average Frontogenesis

Ganetis et al. (2018)

Snow bands and surface snow rates



To what degree do locally- enhanced reflectivity banded features have an impact on surface snow fall rates?

Data: Hourly ASOS surface station data (2012-2023)



٠

NEXRAD stations with 200 km range ring

17,486 total samples over 29 stations and 264 storm days

Only using stations with AWPAG weighing gauges with wind shields for best LWE measurements of frozen precipitation

Data: NEXRAD regional radar maps

45

40

30

25

20



Radar Reflectivity [dBZ] melting and mixed precip regions removed from analysis Tomkins et al. (2022)

Snow Rate [mm hr⁻¹] Rescaled Z to better represent snow field using following equation: $Z_{p} = 57.3 \ S^{1.67}$ Rasmussen et al. (2003)

Snow rate [mm hr⁻¹]

10

8



Objective Feature Detection for snow bands Tomkins et al. (2023)

Faint: not very distinct features **Strong**: very distinct features Background: echo surrounding objects

netCDF files for all regional composites between 1996-2023 available on Dryad (open-source data repository) soon

• 1996-2012 (pre dual-pol) DOI: 10.5061/dryad.zcrjdfnk6

2012-2023 (post dual-pol) DOI: 10.5061/dryad.rbnzs7hj9

Methods: Combining ASOS hourly surface station data and feature detection regional radar maps



Most of the time, locally enhanced reflectivity is associated with low snow rates.



Example: High feature area, heavy snowfall





Example:

High feature area, low/moderate snowfall





Example:

Low feature area, low/moderate snowfall





Example: Low feature area, High snowfall





Big picture takeaways from multi-year winter storm observations

Low correlation between enhanced Z in "snow bands" detected on regional scanning radar and hourly surface snow rates

- Usually observe coincident mixtures of snow particle shapes and degrees of riming
 - Mixtures yield varying distributions of shapes, sizes, and densities in the same volume which complicate interpretations and retrievals of snow rate
- In the 1-2 hours that it takes a precipitation-sized ice particle to fall from near cloud top to the surface, 3d ice streamers originating in generating cells are tilted and smeared
- Lack of vertical column continuity in local enhancements in radar Z

Further Thoughts

- In regions without distinct upward forcing (e.g. strong frontogenesis, orographic lifting), refrain from labeling as "snow bands" most observed mesoscale linear features of higher reflectivity in winter storms
- Beware of conflating *forecast model* predictions of locally enhanced snowfall based on ice water contents that are converted to radar reflectivity *for display* using simple Z-S relations (where there is a 1:1 between increasing snow rate and increasing Z) with *observed* radar reflectivities
- Suggest better bet for evaluating model predictions of snowfall are hourly surface snowfall liquid equivalent measurements (in non-blizzard conditions) rather than retrievals from observed radar data